**Original Research** 

# Q-SSC Behavior During Floods in the Isser Watershed, (North-West of Algeria)

# Djouhra Baloul<sup>1, 2\*</sup>, Abderrahmane Nekkache Ghenim<sup>1</sup>, Abdesselam Megnounif<sup>1</sup>

<sup>1</sup>'Eau et Ouvrage dans Leur Environnement' Laboratory, University of Tlemcen, BP 230, 13000 Tlemcen, Alegria <sup>2</sup>Institute of Technology, University of Bouira, 10000 Bouira, Algeria

> Received: 5 December 2023 Accepted: 5 March 2024

#### Abstract

The study aims to establish a graphical relationship between sediment concentration (C) and water discharge (Q) during flood events in the Isser catchment. Hysteresis, indicating a time lag between discharge flow (Q) and suspended sediment concentration (SSC) curves, varies based on sediment availability, event magnitude, and sequence. Based on the 2026 data pairs of water discharge and suspended sediment concentration (Q-SSC), we have selected 22 flood events. The most frequent hysteresis loops were complex (10 loops), with 08 clockwise loops, 02 figure-eight loops, and 02 anti-clockwise loops. Complex hysteresis loops accounted for 63% of solid loads and 37% of water discharge loads, while 50% of total water yield and 23% of total sediment yield were associated with clockwise loops. Principal Component Analysis (PCA) revealed that water discharge load, mean concentration, maximum concentration, and concentration at the flow discharge peak are key variables influencing hysteresis patterns.

Keywords: sediment concentration, discharge, hysteresis, ACP, Wadi Isser

### Introduction

Estimation, quantification, and comprehension of sediment transportation in dynamic systems are important for project planning: estimation of dam silting, degradation of aquatic habits, coastline changes [1], and developing efficient agricultural land management strategies [2-4]. Erosion, soil loss assessment, water treatment, association of pollutants with sedimentary particles, and particle depositions at dam reservoirs are all outcomes of suspended sediment phenomena [1, 5, 6]. Quantities of sediment drained to the outlet are linked to several complex factors, such as catchment characteristics, land occupation, channel morphology, initial soil moisture, and hydrological and climatic conditions [7-9].

Estimation of sediment concentration is commonly performed using sediment rating curves with a power function. These curves represent the average relationship between discharge and suspended sediment concentration within a specific watershed [10-13]. Analyzing the Q-SSC relation graphs during flood events allows for the identification of sediment source patterns, which sometimes show hysteresis [14].

Hysteresis pattern is a graphical tool used to analyze and understand the behavior of sediment concentration

<sup>\*</sup>e-mail: hydraulique.eau@gmail.com

in relation to discharge (C-Q) at a specific storm event [3, 8, 15-17]. Hysteresis appears when the time lag between the discharge curve and the suspended sediment concentration (SSC) curve is different [18]. Hysteresis is influenced by various factors, including the magnitude and sequence of events, sediment particle size distribution, basin size, land use, and sediment sources. Hysteresis between C-Q was first analyzed by [19]. According to the delay in the peak of the hydrological graph and sediment graph [19], hysteresis is categorized into five common patterns: (I) a single line, (II) clockwise, (III) counterclockwise, (IV) a single line plus a loop, and (V) a shape eight. For a more detailed explanation of the various classes of hysteresis, consult the work of Williams (1989) [20].

Sediment storage along transport pathways was responsible for clockwise hysteresis due to the availability of sediment in streams, leading to the "first flush" effect. Counterclockwise hysteresis appeared to be attributed to enhanced downstream and lateral hydrological connectivity on the surface [21]. The building and operation of water regulation structures modified the SSL-Q supply, carrying capacity, and erosive behavior of the river. These changes can be observed in the sediment rating curve and hysteresis analyses [4].

A hysteresis research has predominantly been conducted in advanced nations. However, there is a pressing need for more extensive investigations, particularly in the Southern Hemisphere, where a distinct opportunity exists to gain insights into hysteresis within diverse climatic and catchment contexts [18]. Moreover, the predominant focus in hysteresis research has been on smaller catchments ( $<10^3$  km<sup>2</sup>), where relationships are more straightforward to interpret. Conversely, the application of SSC-Q hysteresis analysis to larger catchments (>103 km2) poses inherent challenges in comprehending sediment dynamics [4].

The Algerian catchment is characterized by contrasting climatic conditions, with summer wind storms and unpredictable fluctuations in rainfall on

E

an annual and intra-annual scale. Frequent heavy rainfall in the region occurs on arid, sparsely vegetated soil, leading to robust surface runoff, which results in substantial soil erosion [8, 22]. This leads to complex flow and sediment concentration during floods.

The aim of this study is to (1) characterize the SSC hysteresis loops and (2) assess the persistency of hysteresis loops across various timescales in the Isser watershed. The outcomes of this case study may offer insights into suspended sediment transport dynamics and sediment budgets, guiding the adoption of effective soil and water conservation measures. This study was conducted using 16 years of data recorded in the Wadi Isser, which flows into the El-Izdihar dam. This last has an initial volume of 110 Hm<sup>3</sup> (in 1988), with an annual siltation forecast at about 800 000 m<sup>3</sup> [23].

# **Material and Methods**

### Study Site and Data Collection

The Isser catchment is the most significant tributary of Oued Tafna due to its extensive course rather than its strong flow. It is located to the east of the Tafna watershed (Fig. 1). Its drainage area is about 1140 km<sup>2</sup>, length of the main stream is about 81 km. The principals' characteristics of this basin are given in Table 1.

The basin is characterized by a Mediterranean climate; the watershed's climatology indicates an annual average rainfall of 389 mm. In wet years, precipitation can soar to 600 mm, whereas in dry years, it may drop to less than 280 mm (Fig. 2). The average annual water load is 194x10<sup>6</sup> m<sup>3</sup>, while the average annual suspended sediment load is 87×10<sup>3</sup> tons, with a specific degradation of about 77 tons/km²/year.

The southern part of the Isser catchment is dominated by maquis and degraded forest cover, interspersed with some cereal and seasonal crops, which increases erosion. However, the northern part is more land-covered.

20 km



Area A (Km <sup>2</sup> )	1140				
Perimeter P (km)	180.95				
Gravelius compactness coefficient Kc	1.5				
drainage density Dd	1.08				
Roche index slope (%)	0.1				
Overall slope index (m/km)	0.011				
Torrentially coefficient	2.1				
Average concentration time (h)	14 h and 30 min				
Minimum elevation	275				
Maximum elevation	1625				

Table 1. Characteristics of Wadi Isser.

The Isser basin was controlled by the Sidi Aissa hydrometric station (code: 16-06-14) (x = 157.35; y = 199.50; z = 380.0). The instantaneous data measurement of water flow Q and sediment concentration (C) used in this study ranged from 1988-89 to 2003-04 (Fig. 3) was made available by the National Agency of Hydraulic Resources (ANRH). The technique of sampling is the same at the Algerian hydrometric station [17, 24]. The temporal variation of flow Q and sediment concentration C during the study period is illustrated in Fig. 2.

The liquid contribution  $Al_i$  of period (year, season, month and flood event) is given by:

$$Al_{i} = \sum_{0}^{t} \frac{1}{2} (Q_{i} + Q_{i+1})(t_{i+1} - t_{i})$$
(1)

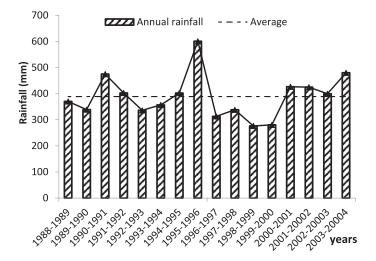


Fig. 2. Rainfall variation.

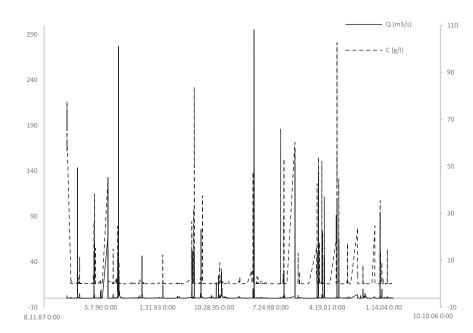


Fig. 3. Variations of instantaneous measurements of water discharge and suspended sediment concentration during 1988 to 2004 at the Sidi Aissa gauging station.

The solid contribution  $As_i$  for a period, is given by the relation:

$$As_{i} = \sum_{0}^{t} \frac{1}{2} (C_{i}Q_{i} + C_{i+1}Q_{i+1})(t_{i+1} - t_{i})$$
(2)

#### Hysteresis Analysis

#### Classification

In the literature, there are five classes of hysteresis:

Class I: occurs when there is similarity in the SSC-Q relation in the hydrograph's rising and filling limbs, there is no hysteresis loop observed, and the sediment concentration and water discharge follow a single line without any significant delay or loop formation.

Class II: the peak of SSC occurs behind the peak of water discharge. In this type, as the water discharge increases, the sediment concentration initially remains low but gradually rises. As the discharge decreases, the sediment concentration remains relatively high before eventually decreasing. This hysteresis pattern often indicates sediment storage and delayed sediment response within the watershed.

Class III: in a counterclockwise hysteresis loop, sediment concentration increases rapidly as water discharge rises, reaching a peak at high water flows. As water discharge decreases, sediment concentration decreases more slowly, resulting in a counterclockwise loop. This type of hysteresis loop typically occurs when sediment transport is dominated by a suspended load and there is a significant supply of fine sediments.

Class IV: the figure-eight hysteresis pattern shows a more complex behavior, characterized by the formation of two distinct loops resembling the shape of a figure- eight. These loops may be clockwise or counterclockwise and indicate different sediment transport behaviors during rising and falling water discharge.

Class V: this shape combines class I with another hysteresis loop. The loop may be clockwise or counterclockwise, indicating different behaviors during rising and falling water discharge. It can contain more than two hysteresis loops, generated by a single or multiple peak flood [9].

Hysteresis was plotted after selecting 22 important flood events. Table 1 summarizes the characteristics of these flood events. To obtain the hysteresis shape, we have plotted for each flood event the hygrograph of flow discharge Q and the hydrograph of suspended sediment concentration C than the plotted graphs C versus Q relation. Fig. 4 illustrates some flood events.

The hysteresis loop offered a visual comparison of pattern, size, and shape. However, researchers used indexes to quantify hysteresis characteristics [18]. Among them:

1. The hysteresis index HI witch quantify the curves or loops at different stage of hydrograph or the area under the curve. It is defined by [9] as:

$$HI = \log\left(\frac{SSC_R}{SSC_F}\right) \tag{3}$$

Where  $SSC_R$  and  $SSC_F$  are the concentration of suspend sediment concentration in rising and filling limb.

2. The peak phase difference P-d-p which is a tool to quantify the manner of the variation of SSC around the peak flow [9] it is given by:

peak phase difference = 
$$\left(\frac{T_Q-max-T_{SSC-max}}{T_{flood}}\right)$$
 (4)

Where  $T_{Q-max}$ ,  $T_{SSC-max}$  and  $T_{flood}$  are respectively the time of the peak of flow, the time of the peak of SSC and the flood duration.

#### **Results and Discussion**

Table 1 groups the characteristics and classes of 22 flood events. The maximum flow discharge recorded was 147.3 m<sup>3</sup>/s, corresponding to the flood event on September 27, 1999. The maximum suspended sediment concentration C observed was 83.76 kg/m<sup>3</sup> during the flood event on November 11, 1994. The minimum flow discharge was recorded during the flood event on January 1, 2003, and the minimum SSC was 1.06 kg/m<sup>3</sup> during the flood event on December 3, 1995.

The longest-lasting flood was approximately 34 days, starting on February 26, 1996, and lasting until March 31, 1996. The shortest flood duration was that of September 27, 1999, lasting for more than 9 hours. The maximum annual contribution of water load was 73.74%, while the maximum annual contribution for suspended sediment concentration C load was 97.4%.

During the study period, 41% of floods occurred in autumn, 36 in winter, 23% in spring, and 0% in summer.

Application of hysteresis techniques for the selected floods highlights 04 different classes (Fig. 5).

Complex loop hysteresis with a single or multiple flow peak was the most dominant hysteresis in the Isser Wadis (45%) contributing to 38% of the total water load and 63% of the total sediment load during the flood event.

The majority of complex loops occur in autumn (September-October-November), the relationship between discharge Q and suspended sediment concentration C is characterized by low water flow and high levels of suspended sediment. A significant portion of the suspended sediments from the wash load coming from the basin is generated by a considerable area.

The flood of 19/3/89 took the form of a complex shape (Fig. 5) with multiple peak flow and precisely multiple peak of SSC, which can be a result of bank erosion after saturation.

The second dominance loop was a clockwise loop with 36% of flood events, 52% of the water contribution,

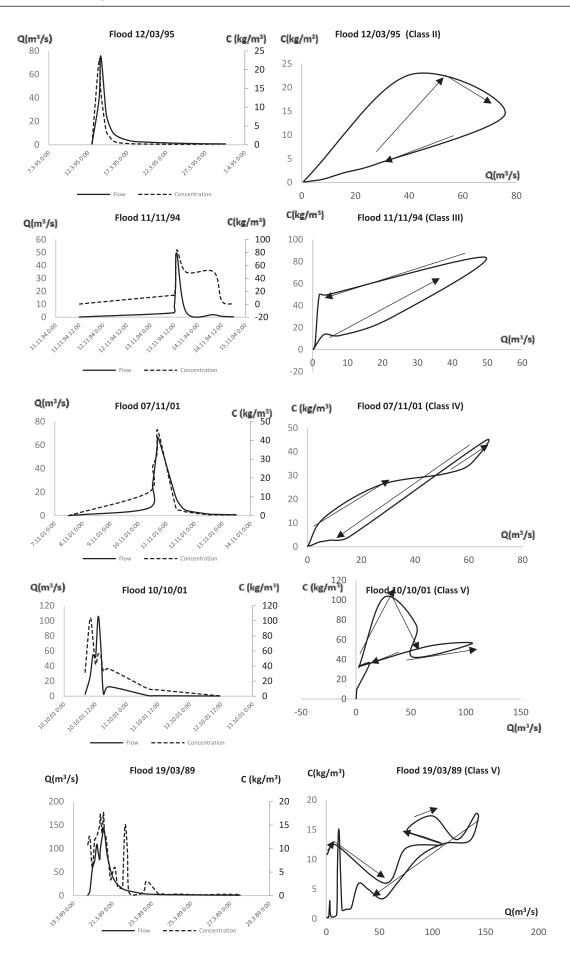


Fig. 4. Example of flood events from different classes.

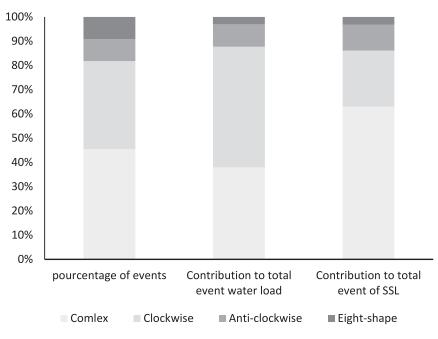


Fig. 5. Contribution of different hysteresis flood events.

and 23% of the solid contribution during flood events. This kind of loop occurs in winter (December-January-February) and at the beginning of spring (March). Increasing flow leads to an increase in sediment concentration, and the peak phase difference is low. The relationships between Q and C are very strong; the majority of sediments are derived from the hydrographic network [22].

Previous studies in Algerian watersheds showed the dominance of clockwise hysteresis; it represented 60% of floods in the Kebir West River [24], 37% of floods in Wadi Sebdou [8], and 50% of floods in the Mekerra catchment [15].

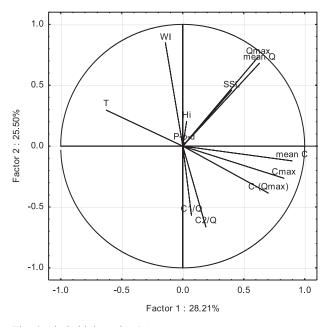


Fig. 6. Circle biplot using PCA test.

Classes III, manifested only by 02 floods at the end of autumn (November) and winter (January), generated 03% of the solid load. This shape occurs when the soils are near saturation. Additionally, it may result from a minor rainfall event following a major one, leading to the erosion of stream banks [22].

The figure-eight (class IV) represented 9% of flood events and 11% of the solid flux. It occurs in January and November.

Fig. 6 shows that factor 1 explains 28.21% of the variance of the origin dataset and factor 2 explains 25.5% of the variance of the origin dataset; therefore, more than 50% of the variance is explained by the first two components. No single line hysteresis was found, and concentration variations do not follow flow rate variations. Watershed geometry can justify the absence of such a form in the Isser basin.

The HI index (Table 2) varied between -0.063 and 2.78; only three values closed to 0, which indicated a small difference in the Q-SSC relationship between the rising and falling stages [10]. For the majority of hysteresis floods, the HI indexes indicated a higher value, which reflected the complex behavior between C and C. All peak phase difference values (Table 2) closed to 0, and this extended between Q and C peaks.

The Principal Component Analysis PCA was employed to explore liaisons between different parameters tabulated in Table 2 and determine which of these parameters governed and influenced hysteresis patterns. According to the circle biplot in Fig. 6, water load was the best variable, with loading close to 0.9 for Factor 2 affecting the hysteresis model.

The mean concentration, maximum concentration, and concentration corresponding to the flow discharge peak have loadings higher than 0.7 for factor 1, which

<b>F1 11</b>	CI	0		G	01/0	02/0	6	~	T.T.	m	D 1 1 1'''
Flood date	Class	Qmax	C-(Qmax)	Cmax	C1/Q	C2/Q	mean Q	mean C	Hi	Т	Peak phase difference
19/03/1989	V	143.4	17.16	17.68	0.120	0.123	44.862	6.9	0.144	7.771	0.0017
04/01/1990	V	28.44	11.62	26.26	0.409	0.923	6.736	7.96	-0.815	3.208	-0.1623
07/09/1990	V	112	43.54	45.28	0.389	0.404	34.74	32.504	0.382	3.139	-0.018
01/03/1991	V	32	9.63	24.39	0.301	0.762	9.644	3.881	0.625	13.125	0.1143
14/03/1991	II	277	20.08	33.77	0.072	0.122	77.311	9.872	1.469	5.417	0.0058
29/09/1994	V	53.82	3.35	26.54	0.062	0.493	9.074	11.187	0.902	4.208	0.307
11/11/1994	III	49.5	83.76	83.76	1.692	1.692	10.02	27.216	0.515	3.25	0
12/03/1995	II	75.7	14.32	22.58	0.189	0.298	17.51	4.451	0.455	11.99	0.0174
03/12/1995	II	17.1	76	1.06	4.444	0.062	4.369	0.386	0.751	8	-0.0104
12/01/1996	III	24.78	2.48	2.48	0.100	0.100	4.595	0.989	0.277	10	0
26/02/1996	II	28.44	6.4	6.4	0.225	0.225	6.822	2.133	0.267	34	0
13/03/1999	II	24.98	53.2	53.2	2.130	2.130	5.414	13.371	-0.063	7	0
27/09/1999	V	147.3	57.68	60.26	0.392	0.409	51.031	44.123	0.172	0.389	0.134
17/10/2000	V	31.48	16.36	42.59	0.520	1.353	5.251	12.683	1.171	9.979	0.008
13/11/2000	V	23.8	22.55	53.92	0.947	2.266	7.233	16.246	0.119	12.024	0
28/01/2001	IV	48.28	21.6	22.66	0.447	0.469	16.016	6.498	1.579	11.01	0.0151
26/02/2001	II	44.82	1.4	37.16	0.031	0.829	8.565	4.267	1.044	14.99	0.08
10/10/2001	V	105.8	56.14	102.6	0.531	0.970	29.541	40.836	0.603	2.146	0.0605
07/11/2001	IV	66.24	44.64	44.64	0.674	0.674	20.173	13.187	2.776	5.979	0
01/01/2003	II	9.92	6.16	7.88	0.621	0.794	3.966	2.743	0.479	3.021	0.03
08/11/2003	V	94.28	35.38	35.38	0.375	0.375	13.016	15.936	0.095	11	0
10/12/2003	II	10	3.68	9.26	0.368	0.926	3.751	3.096	0.923	6	0.1076

Table 2. Characteristics of flood events at Isser wadis.

Qmax, Cmax, Qmean, Cmean: are respectively maximum, mean water flow  $(m^3/s)$  and sediment concentration  $(kg/m^3)$  during flood event, C(Qmax): value of sediment concentration corresponding with Qmax, T: time duration of flood (day).

means they govern the hysteresis shape. Their closed axes reflect a good correlation between these parameters.

These results agree with [18]. For medium and large basins, precipitation, discharge, and sediments are the mean variables that control hysteresis.

The hysteresis pattern is significantly influenced by the mean and maximum flow discharge, which have a good correlation. Peak phase difference P-p-d and HI index were poorly represented, with low loadings for Factor 1 and Factor 2.

# Conclusions

The current study elaborated on the hysteresis method during flood events in the Wadi Isser catchment based on the 2026 data pairs of water discharge and suspended sediment concentration Q-C. Following 22 flood events, we observed four categories of hysteresis patterns, with the complex shape dominating at 36%, followed by the clockwise loop at 23%.

The remaining portion was distributed equally between the anti-clockwise and eight-shape patterns. This classification contributes differently to water and solid loads during flood events. Flood events with clockwise hysteresis dominate in terms of water discharge load, accounting for 50% of the contribution. On the other hand, flood events with complex hysteresis acquire the highest contribution in terms of solid load, amounting to 63%.

The PCA analysis reveals that the parameters controlling the hysteresis pattern in the study area are principally: water load, mean concentration, maximum concentration, and concentration corresponding to the flow discharge peak.

Studying hysteresis in flood events is a crucial tool to improve our understanding of the factors that influence suspended sediment dynamics. Which is important for effective water resource management and sediment control strategies in Algerian watersheds. It helps, especially during storm events, which significantly contribute to the total suspended sediment load. The complex interrelationships among variables such as rainfall intensity, spatial distribution, runoff volume, and flow rate pose a significant challenge to achieving more accurate predictions.

# **Conflict of Interest**

The authors declare no conflict of interest.

#### References

- 1. GRAUSO S., PASANISI F., TEBANO C. Modeling the suspended sediment yield in Lesotho rivers. Modeling Earth Systems and Environment. 6 (2), 759, 2020.
- VALE S.S., FULLER I.C., PROCTER J.N., BASHER L.R., DYMOND J.R. Storm event sediment fingerprinting for temporal and spatial sediment source tracing. Hydrological Processes. 34 (15), 3370, 2020.
- LLOYD C.E., FREER J.E., JOHNES P.J., COLLINS A. Using hysteresis analysis of high-resolution water quality monitoring data, including uncertainty, to infer controls on nutrient and sediment transfer in catchments. Science of the Total Environment. 543, 388, 2016.
- VALE S.S., DYMOND J.R. Interpreting nested storm event suspended sediment-discharge hysteresis relationships at large catchment scales. Hydrological Processes. 34 (2), 420, 2020.
- KISI O., YUKSEL I., DOGAN E. Modelling daily suspended sediment of rivers in Turkey using several datadriven techniques. Hydrological Sciences Journal. 53 (6), 1270, 2008.
- VERCRUYSSE K., GRABOWSKI R.C., HESS T., LEXARTZA-ARTZA I. Linking temporal scales of suspended sediment transport in rivers: towards improving transferability of prediction. Journal of Soils and Sediments. 20 (12), 4144, 2020.
- RAGHUWANSHI N.S., SINGH R., REDDY L.S. Runoff and sediment yield modeling using artificial neural networks: Upper Siwane River, India. Journal of Hydrologic Engineering. 11 (1), 71, 2006.
- 8. MEGNOUNIF A., TERFOUS A., OUILLON S. A graphical method to study suspended sediment dynamics during flood events in the Wadi Sebdou, NW Algeria (1973-2004). Journal of Hydrology. 497, 24, 2013.
- HADDADCHI A., HICKS M. Interpreting event-based suspended sediment concentration and flow hysteresis patterns. Journal of Soils and Sediments. 21 (1), 592, 2021.
- BALOUL D., GHENIM A.N., MEGNOUNIF A. Estimation of Sediment Concentration Using Sediment Rating Curve Approach in Isser Watershed (North-West of Algeria). Ecological Engineering & Environmental Technology. 24 (6), 282, 2023.
- ASSELMAN N.E.M. Fitting and interpretation of sediment rating curves. Journal of Hydrology. 234 (3-4), 228, 2000.
- BOAKYE E., ANYEMEDU F.O.K., DONKOR E.A., QUAYE-BALLARD J.A. Variability of suspended sediment yield in the Pra River Basin, Ghana. Environment, Development and Sustainability. 24 (1), 1258, 2021.

- SADEGHI S.H.R., SAEIDI P., SINGH V.P., TELVARI A.R. How persistent are hysteresis patterns between suspended sediment concentration and discharge at different timescales? Hydrological Sciences Journal. 64 (15), 1909, 2019.
- ELAHCENE O., BOUZNAD I.E., BOULEKNAFT Z. Erosion and solid transport processes in the Isser Wadi watershed, Algeria. Euro-Mediterranean Journal for Environmental Integration. 4 (1), 2019.
- DIAF M., HAZZAB A., YAHIAOUI A., BELKENDIL A. Characterization and frequency analysis of flooding solid flow in semi-arid zone: case of Mekerra catchment in the north-west of Algeria. Applied Water Science. 10 (2), 1, 2020.
- ZOUNEMAT-KERMANI M., MAHDAVI-MEYMAND A., ALIZAMIR M., ADARSH S., YASEEN Z.M. On the complexities of sediment load modeling using integrative machine learning: Application of the great river of Loíza in Puerto Rico. Journal of Hydrology. 585, 124759, 2020.
- BENSELAMA O., HASBAIA M., DJOUKBALA O., BOUTAGHANE H., FERHATI A., DJERBOUAI S. Suspended sedimentary dynamics under Mediterranean semi-arid environment of Wadi El Maleh watershed, Algeria. Modeling Earth Systems and Environment. 8 (1), 1013, 2022.
- MALUTTA S., KOBIYAMA M., CHAFFE P.L.B., BONUMÁ N.B. Hysteresis analysis to quantify and qualify the sediment dynamics: State of the art. Water Science and Technology. 81 (12), 2471, 2020.
- LEOPOLD L.B., MADDOCK JR T. Relation of suspended-sediment concentration to channel scour and fill. Conference Paper, Proceedings of the fifth Hydraulics Conference, Institute of Hydraulic Research, University of Iowa. 1952
- WILLIAMS G.P. Sediment concentration versus water discharge during single hydrologic events in rivers. Journal of Hydrology. 111 (1-4), 89, 1989.
- KEESSTRA S.D., DAVIS J., MASSELINK R.H., CASALÍ J., PEETERS E.T.H. M., DIJKSM R. Coupling hysteresis analysis with sediment and hydrological connectivity in three agricultural catchments in Navarre, Spain. Journal of Soils and Sediments. 19 (3), 1598, 2019.
- 22. MEGNOUNIF A., TERFOUS A., GHENAIM A., POULET J.B. Key processes influencing erosion and sediment transport in a semi-arid Mediterranean area: The Upper Tafna catchment, Algeria. Hydrological Sciences Journal. 52 (6), 1271, 2007.
- ZOBIRI M., MAZOUR M., MORSLI B. Water erosion on marl slopes and prevention of its effects using conservation of water and soil systems in the Wadi Isser watershed -Algeria. Journal of Water and Land Development. 37 (1), 161, 2018.
- 24. KHANCHOUL K., SAAIDIA B., ALTSCHUL R. Variation in Sediment Concentration and Water Discharge During Storm Events in Two Catchments, Northeast of Algeria. Earth Science Malaysia. 2 (2), 01, 2018.
- 25. KHANCHOUL K., SAAIDIA B., ALTSCHU R. Variation in Sediment Concentration and Water Discharge During Storm Events in Two Catchments, Northeast of Algeria. Earth Science Malaysia. 2 (2), 01, 2018.